




Article

GEOCHEMICAL AND ISOTOPE ANOMALIES IN SIOMA RIVER GORGE, CENTRAL TAJIKISTAN

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Abstract: For a better understanding of the previously observed increased natural radioactivity at the confluence of Varzob and Sioma rivers, the activity concentration of natural radioactive ⁴⁰K, Th and U daughters as well as of the anthropogenic ¹³⁷Cs were determined in nine sampling points covering the entire Sioma river. The radiometric determinations evidenced, beside increased activity concentrations of Th and U daughter radionuclides, an anomalous accumulation of anthropic ¹³⁷Cs and natural ²¹⁰Pb in the middle of the Sioma gorge. Complementarily, the mass fractions of natural Th and U determined by Instrumental Neutron Activation Analysis (INAA) evidencing an increased presence of Th and U which, in some places overpasses six times the average mass fraction found in the Upper Continental Crust. The ¹³⁷Cs and ²¹⁰Pb activity concentrations did not correlate with the other radionuclides or Th and U presence, but were reciprocally correlated suggesting an active air transport from the southern Tadjik plane most probably during the Afgan storms. At their turn, the increased presence of Th and U as gamma-ray spectroscopy and INAA determinations proved could be attributed to the existence of actinides rich minerals, similar to those found on neighbouring Odjuk pegmatite field.

Keywords: Sioma river; Tajikistan; natural radioactive elements; radiocesium; pegmatite; gamma ray spectrometry; instrumental neutron activation analysis.)

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1. Introduction

The Sioma River, a right-bank tributary of the Varzob River flows into Varzob about 50 kilometers north of city of Dushanbe, the capital of Tajikistan. The Sioma River Gorge stretches parallel to the Hissar Range having an east-west orientation. The gorge is about 17 km long and rather narrow, its width not exceeding 500 meters. It is framed by mountains with snow and glacial peaks 3500 - 4600 meters high while the gorge altitude varies between 2500-2700 m (Fig. 1).

The basin of the Sioma River, in geological terms, belongs to the South Hissar structural-formational zone of Central Tajikistan (Southern Tien Shan Mountains). Its catchment basin is located in the multi-phase Hissar batholith, composed of various granitoids of the Middle Carboniferous-Early Permian age and broken in places by Early Mesozoic lamprophyre dikes and explosive tubes with alkaline basalts [1]. The climate is cold and temperate, classified as Dfb according to the Köppen-Geiger climate classification [2].

31 Sioma basin is located in an area of aerosol deposition by the Afghani dust storms,
32 but due to the steep banks of the gorge and abundant precipitation, dust particles
33 are washed into the river, although in some places they form shallow loess soils [3,4].
34 During this aerosols transport, first of all, large particles and heavy minerals fall out, the
35 remaining dust is enriched with light minerals, including potassium. At the entrance to
36 the Sioma Gorge, there is a sharp change in the direction of the wind, and it, literally
37 licks the surface of the rock leading to increased fallout of aerosols. At the end of the
38 gorge, winds and dust hit higher rocks, which could also determine an increased aerosol
39 deposition.

40 Previous studies performed within the Cooperative Transboundary Monitoring
41 Data Sharing and Modeling of Water Resources in Central Asia (NAVRUZ) project [5],
42 have evidenced along Varzob river an increased gamma ray activity at the confluence of
43 the Sioma and Varzob rivers [6,7]. This finding was analyzed by taking into account the
44 existence of post Soviet Era of a significant number of former uranium mines mainly
45 located in the northern Tajikistan ridges as well in the other neighbouring mountainous
46 regions of Central Asian states [8–10].

47 In view of these findings, during the summer season of 2019 it was organized by
48 the Institute of Water Problems, Hydropower and Ecology of the Academy of Science of
49 Tajikistan to collect more 9 samples representative for the environmental radioactivity of
50 Sioma river sediments. High resolution gamma ray spectroscopy as well as Instrumental
51 Neutron Activation Analysis (INAA) were used to determine the activity concentration
52 of ^{232}Th and ^{238}U radioactive daughter as well as the mass fractions of natural Th and U
53 in collected sediments.

54 The final aim of this study consists in realizing a detailed description of the natural
55 radioactivity distribution along the Sioma river evidencing the contribution of local
56 petrology as well as the long distance airborne transport, mainly by the Afgani storms
57 [11].

58 The results of this project will be further presented and discussed.

59 2. Materials and Methods

60 2.1. Sampling

61 The position of sampling points are shown in Figure 1 and presented in Table A1.
62 It should be noted that, due to water fast flow and granite bed of the river, sometime
63 it was relatively difficult to collect samples of bottom sediment. At the same time, to
64 increase the sampling accuracy, the distances between collecting points were chosen as
65 possible as equal and spread on both river sides.

66 Therefore, nine surface sediments samples were collected by means of an AMS 12
67 Multi Stage Sediment Sampler (AMS Inc. American Falls, ID) [12]. At each sampling
68 point, three independent samples were collected from an area of about 10 m^2 and mixed
69 together which resulted in an average specimen. To avoid any cross contamination,
70 after each operation, sampling tubes were intensively cleaned with water and soaked
71 with paper napkins. Collected sediments of about 1 kg consisting of a silty greyish sand
72 without visible vegetal debris, were kept in clean polyethylene bags and transported
73 to Dushanbe laboratory. Here, the large fragments of rocks and pebbles were removed,
74 the remaining sedimentary material being air dried, sieved through 18 mesh sieve and
75 stored for future analysis [7,13].

76 2.2. Radiometric Measurements

77 Sediment samples, prepared as mentioned before, were poured into 500 cm^3
78 Marinelli backers, sealed and kept four weeks to reach the radioactive equilibrium.
79 Measuring time was of 6 hours for sediment and 24 hours for the background.

80 High resolution gamma-ray spectrometry was used to determine the activity concen-
81 tration of anthropogenic ^{137}Cs , natural ^{40}K as well as ^{212}Bi and ^{212}Pb (^{232}Th daughters)
82 and ^{214}Bi , ^{214}Pb and ^{210}Pb (^{238}U daughters) Al measurements were done using an Ex-

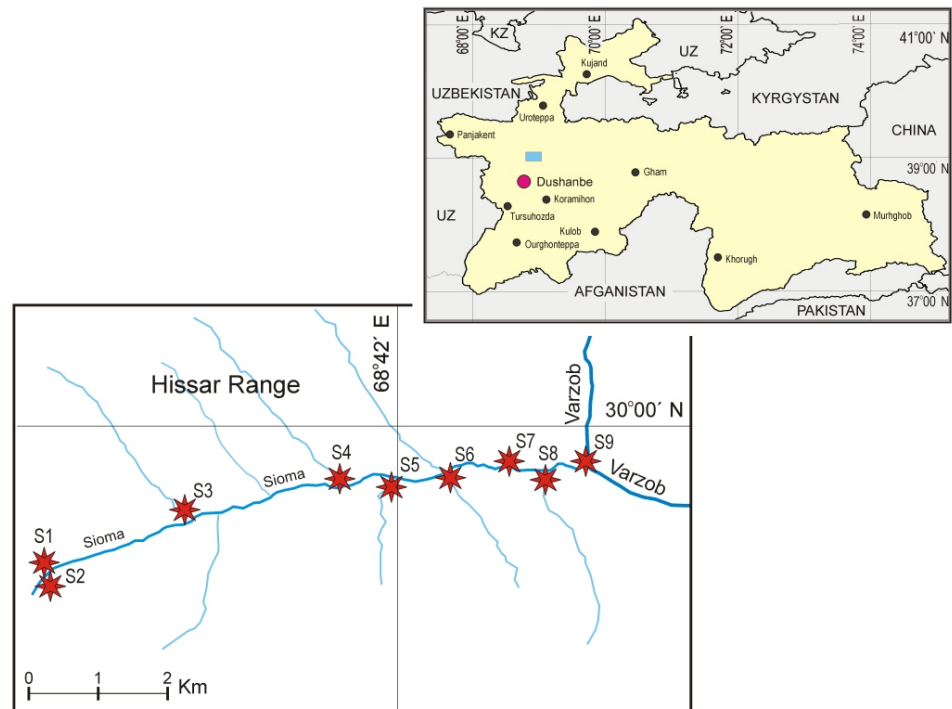


Figure 1. The locations of sampling points along the Sioma gorge whose position with respect to Tajikistan is illustrated in the inset. The tributaries of the Sioma river are in fact small creeks, majority without names.

83 tended Range Coaxial Ge Detector (XtRa) with an energy resolution of 1.78 keV for the
 84 ^{60}Co 1.33 MeV line. More details concerning efficiency calibration can be found in [7].
 85 Under these circumstances, the total uncertainty concerning the activity concentrations
 86 was no greater than 7%

87 2.3. Instrumental Neutron Activation Analysis

88 The mass fractions of Th and U in considered sediments were determined by
 89 INAA carried out at the research nuclear reactor of the Institute of Nuclear Physics
 90 of the Academy of Sciences of the Republic of Uzbekistan [14]. Sediments samples
 91 and reference materials weighting about 0.5 g were wrapped in polyethylene bags and
 92 irradiated twice at a thermal neutrons fluency density of $5 \cdot 10^{13}$ neutrons $\text{cm}^{-2}\text{s}^{-1}$; for
 93 15 s for short living radionuclides and after two weeks for 15 h in the same conditions,
 94 for the long living ones.

95 Spectra calibration was done by using internal reference material by micro-pipetting
 96 a determined volume of chloride or nitrate of considered elements on an ashless filter
 97 paper. Zn was used as secondary reference element. The quality control of all determi-
 98 nations was done by measuring in the same condition the certified reference material
 99 IAEA-336 (Trace and minor elements in lichen). In these conditions, the combined
 100 uncertainty concerning Th and U mass fractions was no greater than 6%. More details
 101 concerning INAA measurements can be found in [6].

102 3. Results

103 Final data concerning the experimentally determined activity concentrations as
 104 well as INAA determined mass fractions of Th and U are reproduced in Table A2, and,
 105 illustrated as stacked columns in Fig. 2a (activity concentration) and Fig. 2b (Th and U
 106 mass fractions). In the last case, we have reproduced for comparison the mass fractions
 107 of Th and U corresponding to the Upper Continental Crust (UCC) as provided in [15].

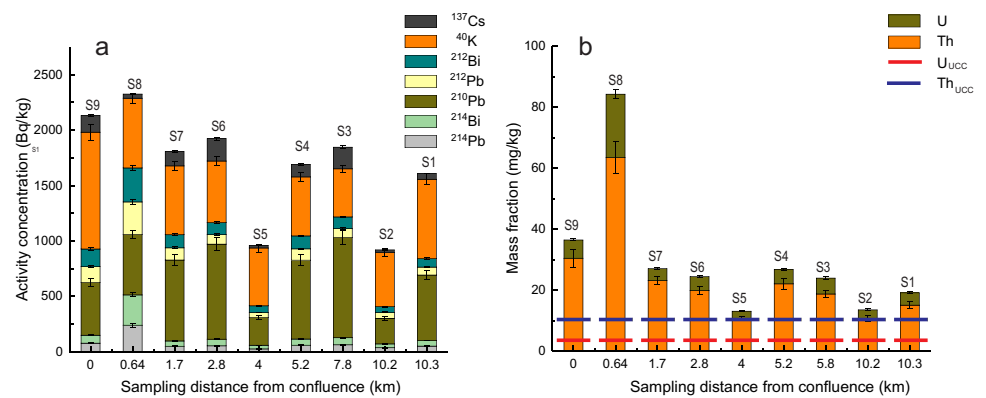


Figure 2. The activity concentration distribution of the main radioactive natural as well as anthropogenic isotopes (a) and the mass fractions of Th and U (b) along the Sioma valley from the confluence with the Varzob river. For comparisons, on fig 2b, the UCC [1] Th and U mass fractions are illustrated too.

108 In the last case, it can be remarked that both Th and U reached maximum mass
 109 fractions for the S8 sampling point slowly decreasing upstream, Th mass fraction signifi-
 110 cantly overpassing with two exceptions, S2 ad S5 sampling points, the UCC one [15]. At
 111 the same time, the Th/U mass fraction ratio varied more or less closer to the UCC one of
 112 3.9 (Table A2).

113 4. Discussion

114 At a more carefully analysis of experimental results illustrated in Fig. 2 and Table
 115 A2, it can be remarked that, the ^{40}K activity concentration presents a relative maximum
 116 at the western end of Sioma Gorge reaching a maximum value of 1060 ± 70 Bq/kg,
 117 slightly higher then the UCC average value of 870 Bg/kg. This value correspond to the
 118 mass fraction of 2.80 % w specific to UCC [15].

119 It worth mentioning that the ^{40}K activity concentration presented two peaks, one
 120 at the entrance in the gorge (S1) and the other one at the confluence with the Varzob
 121 river (S9 on Fig. 1). This behaviour could be associated with the fallout of dusty aerosols.
 122 Indeed, the dust storms known as Afghans [3] originate in the Sahara Desert and fly
 123 a long way through the Arabian Peninsula, Iran and Afghanistan. The approximate
 124 equality of the ^{40}K isotope in the samples from the middle part of the gorge indicates
 125 approximately the same geological and geochemical origin of the soils throughout the
 126 gorge. These are, as mentioned before, loess soils of shallow depth [4].

127 Another evidenced peculiarities concern the mass fraction as well as the activity
 128 concentration of Th and U in conjunction with their daughter radionuclide ^{212}Bi , ^{212}Pb
 129 as well as ^{212}Bi and ^{212}Pb respectively. According to Figure 2a and Table A2, it should be
 130 noted their elevated presence for S8 sampling point. A corresponding sharp increase of
 131 Th and U mass fractions, as illustrated by both INAA and radiometric measurements
 132 (Fig. 2b), excludes, at this point, the possibility of an accidental statistical outlier. To this
 133 remark it should add the fact that ^{212}Pb activity concentration significantly overpasses,
 134 for entire Sioma valley sediments, those of the ^{212}Bi and ^{212}Pb daughters of natural
 135 Uranium (Table A2).

136 On the other hand, ^{212}Pb activity concentration dose not correlates neither with Th
 137 and U daughters radioactivity, nor with the ^{40}K , but only with the anthropogenic ^{137}Cs
 138 one (Table 3a,b). A similar behaviour of the ^{212}Pb activity concentration was previously
 139 reported for the entire length of the Varzob River gorge [7].

140 In spite of this, ^{212}Pb activity concentration presented a maximum at sampling point
 141 S8 where the U mass fraction showed a maximum. This somehow contradictory finding

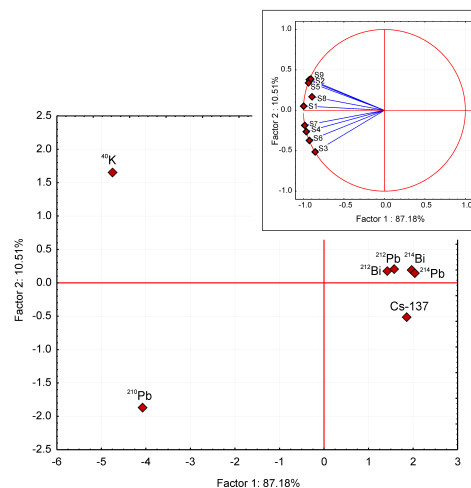


Figure 3. The result of PCA of the activity concentration of investigated radionuclides. While ^{232}Th and ^{238}U daughter, excepting ^{210}Pb form the same cluster proving a correlated distribution of natural Th and U, ^{40}K , ^{137}Cs and ^{210}Pb form different individual clusters suggesting either different parent rocks as the case of ^{40}K or an external source as ^{137}Cs and partially ^{210}Pb .

142 could be explained by an active air transport of radon from the Odjuk pegmatite field as
 143 well as from the entire Hissar granite batholith superposed on the local U sources.

144 Concerning ^{137}Cs , its average activity concentration along Sioma valley of 100Bq/kg
 145 is significantly higher than the average value of 2Bq/kg reported for the entire Central
 146 Asia, most probably the reminiscence of Lob Nor atmospheric nuclear tests [16].

147 Despite different origin and source, between the activity concentrations of both
 148 ^{137}Cs and ^{210}Pb there is a good correlation characterized by a Pearson's correlation
 149 coefficient of 0.824 and a corresponding nonparametric Spearman's correlation coefficient
 150 of 0.812, but at $p < 0.05$ (Table 1). At the same time, between both ^{137}Cs and ^{210}Pb activity
 151 concentrations and the other investigated radionuclide there are no correlations (Table
 152 1).

153 This fact is well illustrated by the Principal Component Analysis (PCA) whose result,
 154 expressed by a Principal Component (PC) 2 vs PC1 bi-plot documents the existence
 155 of at least four clusters, consisting of ^{40}K , ^{137}Cs , ^{210}Pb and ^{212}Bi , ^{212}Pb , ^{212}Bi and ^{212}Pb
 156 respectively (Figure 3).

157 In this regard it is of interest to remark that, by analysing the Factor 2 vs. Factor 1
 158 bi-plot reproduced in Figure 3, inset, results that the major contribution to discrimination
 159 between radionuclides comes, in a more or less equal measure, from all sampling points,
 160 Factor 1 being responsible for more than 81 % of total variance.

161 The increased presence of Th and U indicates the existence of some enriched natural
 162 actinides deposits or ore occurrences which are located on the Sioma gorge in the
 163 vicinity of sampling point S-8, deposits which are either close to the surface or of high
 164 thickness, in good correlation with the previous reported data [7,13] concerning natural
 165 radioactivity along Varzob river

166 The fact that everywhere along the Sioma river, the Th/U ratio varied around 4.3
 167 ± 0.8 , *i.e.*, not far from the 3.9 characteristic of the UCC [15], suggests the presence of the
 168 same type of acid rocks such as granitoids or pegmatites, the last one found, as mentioned
 169 before, in neighbouring Odjuk pegmatite field [7,13]. This observation is sustained by
 170 the evidenced good correlation between Th and U mass fractions characterized by both
 171 Pearson's and Spearman's correlation coefficients of 0.969 and respectively of 0.778 at p
 172 < 0.05 .

Table 1. The matrix of Pearson's (lower diagonal) and Spearman's (upper diagonal) correlation coefficients of activity concentration distribution of investigated radionuclides. The correlations valid at $p < 0.05$ are represented in red ink.

	^{214}Pb	^{214}Bi	^{210}Pb	^{212}Pb	^{212}Bi	^{40}K	^{137}Cs
^{214}Pb	—	0.983	0.307	0.783	0.767	0.383	0.450
^{214}Bi	0.998	—	0.305	0.767	0.750	0.400	0.517
^{210}Pb	0.065	0.048	—	0.250	0.267	-0.117	0.817
^{212}Pb	0.970	0.959	0.099	—	0.983	0.600	0.383
^{212}Bi	0.966	0.955	0.163	0.994	—	0.617	0.400
^{40}K	0.152	0.111	-0.159	0.304	0.293	—	0.133
^{137}Cs	-0.175	-0.203	0.824	-0.086	-0.009	0.090	—

173 5. Concluding Remarks

174 Gamma ray spectrometry was used to determine the distribution of specific activity
 175 of Th and U daughter radionuclides ^{212}Bi , ^{212}Pb , and ^{212}Bi , ^{212}Pb as well as ^{210}Pb respec-
 176 tively together with ^{40}K and anthropogenic ^{137}Cs along the Sioma river, a high mountain
 177 tributary of the Varzob river. These results were analysed in correlation with the Th and
 178 U mass fraction distribution determined by INAA for the same sampling points.

179 The final results showed an increased presence of Th, U and their daughter radionu-
 180 clides near the confluence of Sioma and Varzob rivers attesting the presence of some
 181 mineral fractions, most probable pegmatite similar to those existing not far away on the
 182 Odjuk river valley.

183 Quite different was the situation of ^{137}Cs and ^{210}Pb whose content was significantly
 184 higher than those reported outside the mountainous areas of Tajikistan, which, most
 185 probable, are transported by the local Afgan storms and deposited when the air mass
 186 intercept the Hissar range whose, east-westward orientation facilitate the aerosol deposi-
 187 tion.

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 189 O.G.D.; software, Dj.A.A., O.G.D., and D.A.; validation, Dj.A.A., and O.G.D.; formal analysis,
 190 Dj.A.A., and M.V.F.; investigation, Dj.A.A., and D.A.; resources, Dj.A.A.; data curation, Dj.A.A.,
 191 and O.G.D.; writing—original draft preparation, Dj.A.A., and M.V.F.; writing—review and editing,
 192 Dj.A.A., O.G.D., and M.V.F.; visualization, M.V.F., and Dj.A.A.; supervision, Dj.A.A. and F.M.A.;
 193 project administration, Dj.A.A.; funding acquisition, Dj.A.A.; All authors have read and agreed to
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203 **Appendix A****Table A1.** Coordinates of sampling points, altitude above sea level (m) and distance from the entrance to Sioma gorge - confluence with Varzob river (km)).

Sampling point	Latitude	Longitude	Altitude	Distance	Position
S1	38.9668	68.7503	2631	10.3	Left bank
S2	38.9668	68.7422	2705	10.2	Right bank
S3	38.9651	68.7284	2478	7.8	Left bank
S4	38.9638	68.7154	2184	5.2	Right-bank
S5	38.9642	68.7014	2123	4.0	Left-bank
S6	38.9561	68.6716	2057	2.8	Left-bank
S7	38.9444	68.6482	1893	1.7	Left-bank
S8	38.9478	68.6467	1871	0.6	Right-bank
S9	38.9683	68.7596	1751	0	Varzob confluence

Table A2. The distribution of activity concentration of investigated radionuclides (in Bq/kg) as well as the corresponding mass fractions of Th and U (in mg/kg).

Sampling point	⁴⁰ K	¹³⁷ Cs	²¹⁴ Pb	²¹⁴ Bi	²¹⁰ Pb	²¹² Pb	²¹² Bi	Th
S1	710±50	53±3	51±3	51±3	590±38	71±5	78±7	15.1±1.1
S2	490±35	21±2	37±3	35±3	230±9	54±3	52±4	10.7±0.9
S3	430±30	195±13	64±4	64±4	900±55	80±8	106±8	18.8±1.2
S4	530±35	113±7	59±4	56±4	710±42	102±7	117±10	22.1±1.8
S5	520±30	24±2	27±2	32±2	250±15	45±3	61±6	10.7±0.7
S6	550±40	200±12	55±4	59±4	860±55	88±4	109±8	19.9±1.3
S7	620±40	130±7	49±3	49±3	730±35	111±9	118±10	23.2±1.4
S8	630±40	37±3	240±15	274±19	550±30	293±21	306±25	63.5±5.3
S9	1060±70	150±9	77±4	72±5	480±30	143±11	159±11	30.5±3.0

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